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ANALYSIS AND PREDICTION OF OUTRUNNING GROUND MOTION

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23 January 1985



Technical Report

CONTRACT No. DNA 001-84-C-0125

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110 A/66/12

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13 REPORT SECURITY CLASSIFICATION UNCLASSIFIED		16 RESTRICTIVE	MARKINGS		
2a SECURITY CLASSIFICATION AUTHORITY N/A since Unclassified		3 DISTRIBUTION AVAILABILITY OF REPORT			
2b DECLASSIFICATION/DOWNGRADING SCHEDULE N/A since Unclassified		Approved for	or public re on is unlimi	lease;	
4 PERFORMING ORGANIZATION REPORT NUMBE	R(S)	5 MONITORING	ORGANIZATION RE	PORT NU	MBER(S)
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6a NAME OF PERFORMING ORGANIZATION Applied Research Associates,	6b OFFICE SYMBOL (If applicable)	7a NAME OF MO Director	ONITORING ORGAN	IZATION	
Inc.	(applications)		clear Agency		
6c ADDRESS (City, State, and ZIP Code)			y, State, and ZIP C		
4300 San Mateo Blvd. NE, Suite	A220	Washington	, DC 20305-	1000	
Albuquerque, NM 87110					
8a NAME OF FUNDING SPONSORING	8b OFFICE SYMBOL	9 PROCUREMENT	T INSTRUMENT IDE	NTIFICAT	ION NUMBER
ORGANIZATION (If applicable)		DNA 001-84	-C-0125		
8c ADDRESS (City, State, and ZIP Code)	<u></u>	10 SOURCE OF F	UNDING NUMBERS	5	
	•	PROGRAM ELEMENT NO	PROJECT NO	TASK NO	WORK UNIT ACCESSION NO
		62715H	Y99QMXS	В	DH008148
11 TITLE (Include Security Classification)		3272011	1334		10.000110
ANALYSIS AND PREDICTION OF OUTR	UNNING GROUND MO	OTION			
12 PERSONAL AUTHOR(S)			 		
Hassiotis, Sophia					
Technical Report 13b TIME CO FROM 83	1208 to 850123	850123	RT (Year, Month, E	Day) 15	PAGE COUNT 52
16 SUPPLEMENTARY NOTATION This work was sponsored by the	Defense Nuclear	Agency unde	r DULTE DWCC	Code	
B344084466 Y99QMXSB00008 H259QD).	Agency unde	I NOTOL NISS	code	
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FIELD GROUP SUB GROUP	- Outrunning		Prediction		
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concentrate on a limited number	of HE experime	nts. This r	eport descri	bes th	e development
of a new method to predict the	outrunning port	ion of the q	round shock.	This	method is an
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PREFACE

This report summarizes the work performed by Applied Research Associates on Task 3 of Contract DNA 001-84-C-0125. This task involved the analysis and modification of the prediction procedures for outrunning ground shock.

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CONVERSION TABLE

PERTINENT CONVERSION FACTORS--SI to U.S. CUSTOMARY UNITS OF MEASUREMENT

<u>Parameter</u>	To Convert From	<u>To</u>	<u>Multiply By</u>
Length	Meters (m)	Feet (ft)	3.28
Velocity	Meter/Second (m/s)	Feet/Second (ft/s) 3.28
Density	Kilograms/Cubic Meter (kg/m ³)	Pounds/Cubic Feet (pcf)	62.4 x 10-3
Pressure or Stress	Mega Pascals (MPa)	Pounds/Square Inc (psi)	h 145
Impulse	Mega Pascals Seconds (MPa-s)	Pounds/Square Inc (psi-s)	h-s 145

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SECTION 1 INTRODUCTION

One of the major recommendations of the President's Commission on the Modernization of Strategic Forces (Ref. 1) was the development of a land-mobile, small missile. An outgrowth of this recommendation was the hard mobile launcher concept. Because this is a low overpressure threat, the need for an accurate prediction procedure for the outrunning ground shock wave has increased. The outrunning wave, as defined in this report, is the low frequency ground motion that arrives before the airblast. This wave may be due to the energy coupled at ground zero or to the upstream airblast. The objective of this report is to evaluate the accuracy of existing prediction methods for the outrunning waveform and, if necessary, propose a new prediction method.

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In the past, the prediction of ground motions induced by High Explosive (HE) and Nuclear Explosive (NE) experiments has been the subject of extensive research. Such ground motions can be characterized as follows: 1) airblast-induced, 2) direct-induced, or 3) upstream-induced. Airblast-induced motions are due to the local airblast at the point of interest. Direct-induced motions result from energy coupled at ground zero. Upstream-induced motions are caused by the upstream airblast.

Depending upon explosive yield, height-of-burst (HOB), and geology, ground disturbances may arrive at the point of interest prior to the airblast (i.e., "outrun" the airblast). These disturbances are direct-induced or upstream-induced in nature. At low yields and deep surface layers, outrunning occurs when the airblast velocity falls below the compressional wave speed of the surface layer. At higher yields, or shallow surface layers, faster underlying layers transmit the ground shock. In these cases, outrunning can occur at relatively high overpressures.

Outrunning in the first layer is shown in Figure 1. As seen in this figure, the outrunning region is subdivided into the transeismic and subseismic regions. In the transeismic region, the P-wave velocity is greater than the airblast velocity, which in turn exceeds the S-wave velocity, i.e.,

$$C_s < U < C_p$$

where:

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Cp = P-wave velocity
U = airblast velocity
Cs = S-wave velocity

In the subseismic region, the airblast velocity is less than the S-wave velocity also, i.e.,

Outrunning due to the refraction of energy through a faster, underlying zone is shown in Figure 2. Generally, the refracted P-wave generated by the critically incident P-wave travels parallel to the layer interface at the velocity of the lower layer. This wave is the source of head waves which propagate into the surface layer. Head waves due to the energy at the origin are the primary source of the outrunning seen in data.

Several methods exist for the prediction of the ground shock induced by explosions. Airblast-induced ground motions are predicted with reasonable confidence using finite difference and finite element computer codes. Upstream-induced ground motions, which arrive after the airblast, can be predicted using semi-empirical methods (Ref. 2). Several attempts have been made in the past for the prediction of the outrunning motions. The objective of this report is to investigate the methods produced by such attempts and propose a new method for the prediction of the outrunning wave.

The existing outrunning models are investigated in Section II of this report. The development of the new outrunning prediction procedure will be discussed in Section 3. Section 4 lists the conclusions reached in this work.

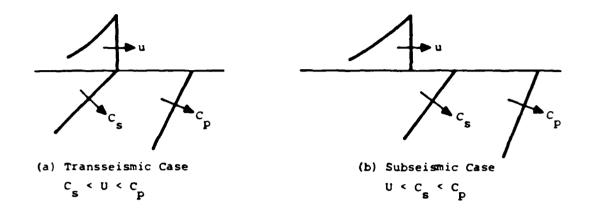


Figure 1. Surface layer outrunning.

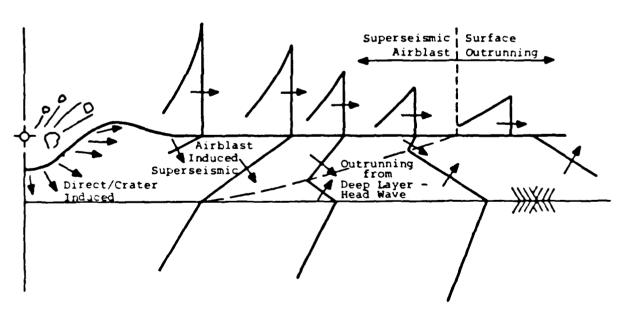


Figure 2. Outrunning due to layer refraction.

SECTION 2 PREVIOUS OUTRUNNING MODELS

Research has been conducted in the past for the theoretical or empirical prediction of the outrunning wave. Four of the methods produced are discussed below.

1) A well-documented method for predicting the entire waveform in the outrunning region (outrunning and upstream-induced that arrive after the airblast) was developed by Sauer (Ref. 3). In this method, a characteristic waveform for the vertical motion was developed using ground motion data obtained from about a dozen surface and above-ground nuclear tests. The data indicated that the period of oscillation increases with range. Thus, a characteristic time, T_2 , was introduced as a scaling factor to time. T_2 is a function of range and is defined as follows:

 T_2 (msec) = 100 + 0.82 ΔR (m) (1) where:

 ΔR = difference between the range of interest and the range at which the first outrunning occurs.

The vertical velocity was normalized with respect to the peak velocity of the waveform (Fig. 3).

Plots from available data provided by DIRECT COURSE, MIXED COMPANY, MILL RACE and other experiments that will be discussed in the next section, indicated that the scaling factors used by Sauer do not collapse the HE data when very different HOBs, yields, and geologies are used. For example, in Figure 4 the curve developed by Sauer is shown with the data from DIRECT COURSE and MIXED COMPANY at the ranges of 466 m and 197m respectively. (The data shown includes only the outrunning portion of the waveform.) As seen, the magnitudes of the outrunning wave are reasonably approximated on the first cycle and overestimated on the second. In addition, the frequency of the MIXED COMPANY wave has been overestimated.

2) A thorough review on the work conducted in the past for the prediction of the outrunning ground motion can be found in Higgins, et al (Ref. 4). Most of this work includes the later time motions (motions that occur after the airblast). A summary

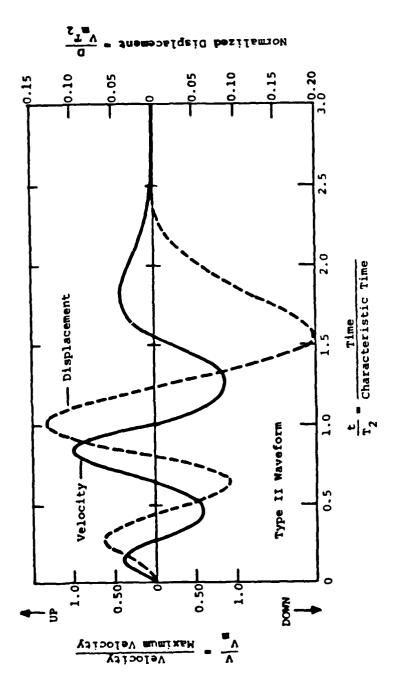
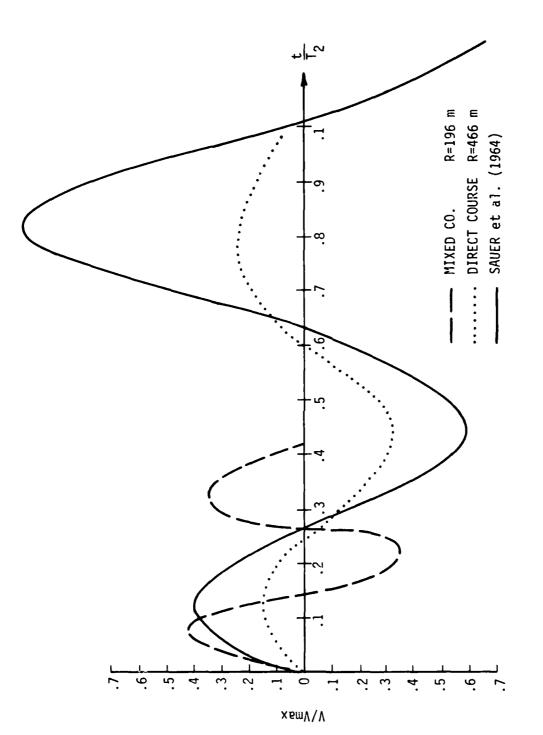


Figure 3. Outrunning waveform. (after Sauer et. al. 1964)



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DIRECT COURSE and MIXED COMPANY data compared to the Sauer waveform. Figure 4.

of this work shows the different attempts made in the past toward an understanding of the motions in the outrunning region. A review of several closed form solutions, numerical calculations, and empirical analyses, and the conclusions reached by such works, are included and briefly discussed in this reference.

The empirical analysis developed in Reference 4 involved a characteristic wave of the entire waveform at the outrunning region, which was developed with data provided by MIDDLE GUST II and III and MIXED COMPANY 3. The velocity was normalized by a value $V_{\rm V}$, equal to the peak velocity of the upstream-induced waveform in the outrunning region. The time was scaled by a value τ that is defined as follows:

$$\tau = \frac{t - t_0}{t_V - t_0}$$
 for vertical
$$\tau = \frac{t - t_0}{t_h - t_0}$$
 for horizontal

where:

t = time after detonation

 t_0 = time of arrival

 t_v = time of occurrence of the third upward peak

 t_h = time of occurence of the first upward peak

The scaling factor

$$\overline{\tau} = \frac{tC_p}{V_a^{1/3}} \tag{3}$$

where:

t = time

 $C_{\rm p}$ = P-wave velocity at a depth below 7.5 m

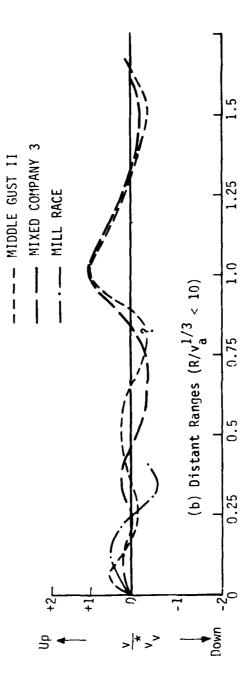
V_a = crater volume

brings the MIDDLE GUST and MIXED COMPANY horizontal velocity data into good agreement.

Figure 5 shows the data of MIDDLE GUST II, MIXED COMPANY 3 and MILL RACE at distant ranges. As seen in this figure, the frequency from the experiments used to develop the method is closely predicted, while the frequency of MILL RACE is underpredicted. In addition, HOB cases can not be included in this analysis, because a crater volume is used in the scaling.

- Another method for predicting the total waveform at the outrunning regions is given by Guros (Ref. 5). By analyzing three 100-ton-HE experiments, Guros derived a characteristic waveform, amplitudes, relations for attenuation with range, time of arrival and frequencies. He also gave directions for the prediction of outrunning in dry and wet site geologies. As in the previous reference, this report also refers to the entire waveform and does not concentrate on the outrunning part of it.
- 4) Finally, the method for the prediction of upstream-induced motions developed by Labreche et al. (Ref. 2) provides an estimate of outrunning motions. The upstream induced motions are made up of two components. These are: 1) the body wave, which travels at P-wave speeds, and 2) the surface wave, which travels at S-wave speeds. The arrival of both components is calculated from critical refraction theory using the source as the origin. The frequency of the body wave motions is governed by the crater volume and the surface wave frequency is governed by the shear properties of the site profile. In general, this method accurately predicts the arrival time of the outrunning motions. The frequency and magnitude, however, are not consistently predicted. This method was developed primarily for use in the superseismic region and therefore was never thoroughly checked in the outrunning region. Furthermore, it is limited to surface bursts and therefore will not allow variations in burst height.

Most of the references cited above concentrate on a limited number of HE experiments, with the exception of the procedure introduced by Labreche et al. (Ref. 2), which was developed primarily for use in the superseismic region, and the procedure introduced by Sauer et al. (Ref. 3), which was



Fits to normalized vertical velocities for various events. (after Higgins et al. 1975) Figure 5.

developed for use in the outrunning region. This second procedure will be used in this report and will be modified using HE data for the prediction of the outrunning portion of the velocity waveform. In Section 3, the development of new scaling factors used in addition to those proposed by Sauer will be presented. Finally, the characteristic waveform shown in Figure 3 will be modified using recent HE data. The waveform produced is general enough to be used for predicting outrunning motions of HE and NE explosions that vary in burst height, geology, and yield.

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SECTION 3 DEVELOPMENT OF A CHARACTERISTIC WAVEFORM

The prediction procedure for the outrunning wave proposed by Sauer et al. (Ref. 3) has been chosen to be modified in this report in order to collapse a wide range of data. For this purpose, additional scaling factors were derived for use in conjunction with those proposed by Sauer. These factors are the velocity, V_1 , which is the peak velocity of the first cycle of motion in the outrunning region, and an impedance ratio, $[(\rho C_p) \operatorname{rock}/(\rho C_p) \operatorname{soil}]$, of the significant layers in the geology. V_1 was used instead of V_{max} proposed by Sauer, while the impedance ratio was used to multiply the factor t/T_2 (Fig. 3). In this section, the derivation of these factors as well as the data on which they are based are discussed.

Several experiments provide a substantial data base on the outrunning region. Good quality data is provided by DIRECT COURSE (Ref. 6), MILL RACE (Ref. 7), and MIXED COMPANY (Ref. 8). Other sources of limited data are Pre-DIRECT COURSE (Ref. 9), MIDDLE GUST II (Ref. 10), III (Ref. 11), and IV (Ref. 12), MISERS BLUFF (Ref. 13), and Pre-DICE THROW (Ref. 14). This data base includes surface tangent and height of burst explosions, as well as a variety of geologies and yields.

Major emphasis in the empirical analysis of these data was placed on the velocity waveform, because the velocity is more easily interpreted than other motion components. The frequency and magnitude of the waveform were the two major parameters considered. The airblast-induced motion or any motion arriving after the airblast was not considered in this analysis.

3.1. Test Site Geologies

DIRECT COURSE was conducted with 600 T of ANFO (equivalent to 500 T of TNT) exploded at an HOB of 50.6 m. The site geology consisted of clayey sand and sandy clay. A hard caliche layer existed at a depth of 10.6 m. MILL RACE was a similar experiment exploded tangent to the surface on the same geology as DIRECT COURSE. The seismic data of this profile, found in Reference 15, are given in Table 1.

MIXED COMPANY was an experiment conducted with 500 T of TNT exploded tangent to the surface. The site was composed of sandy, clayey silt overlying an intermixed formation of siltstone, sandstone, mudstone and

Table 1. DIRECT COURSE and MILL RACE Seismic Data (Ref. 15)

Н	С _р	c _s	Υ
(m)	(m/s)	(m/s)	(kg/m ³)
0 - 0.8	365	245	1600
.8 - 3.5	710	545	1696
3.5 - 10.6	900	494	1696
10.6 - 14.6	1325	640	1712
> 14.6	903	470	1792

Table 2. MIXED COMPANY Seismic Data (Ref. 16)

H	C _p	C _s	Y
(m)	(m/s)		(kg/m ³)
0 - 3	655	393	1920
3 - 122	2286	1493	2160

conglomerates. The bedrock consisted of a massive, cross-bedded sandstone. The seismic data of this profile are given in Table 2 (Ref. 16).

MIDDLE GUST II and III were two experiments, each composed of 100 T of TNT, exploded at an HOB of 4.8 m and at surface tangent, respectively. The site was formed by a layer of alluvial sandy clay overlying Cretaceous Pierre shale. A perched water table existed at about 1.2 m. Table 3 contains the seismic data of this site (Ref. 16).

MIDDLE GUST IV consisted of 100 T of TNT detonated at surface tangent. The site overburden was made up of a very thin layer of clayey top soil overlying a layer of silty clay shale and finally a competent clay shale bedrock. This site had a deep water table, unlike that of MIDDLE GUST II and III. The seismic information is contained in Table 4 (Ref. 16).

MISERS BLUFF was an experiment consisting of 120 T of ANFO (equivalent to 100 T of TNT) exploded at surface tangent. The site profile was composed of wind blown silt, clayey silt, and beds of alternating sand and sand gravel. The bedrock consisted of conglomerate sandstone. The seismic data of this site are found in Table 5 (Ref. 16).

Pre-DICE THROW was an experiment consisting of 100 T of TNT exploded at surface tangent. The site geology had layers of silty clay, fine and course silt, and sand with gravel. The seismic data are found on Table 6 (Ref. 16).

3.2. Time of Arrival

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The range (or time) at which outrunning first occurs can be found using the airblast arrival time, and the standard seismic velocities and layer depths of the site (Ref. 17). Figure 6 shows the relationship between the times of arrival of the upstream-induced or direct-induced ground motion and the airblast. The point at which the two curves intercept is the range (or time) at which outrunning first occurs.

As seen in this graph, the first outrunning is due to upstream airblast. However, observations from the data indicate that the time of arrival of the first outrunning coincides with the arrival of the critically refracted P-wave from the origin. This apparent inconsistency

Table 3. MIDDLE GUST II and III Seismic Data (Ref. 16)

Н	С _р	Cs	Y
(m)	(m/s)	(m/s)	(kg/m ³)
0 - 2.13 2.13 - 4.57	610 1768	122 183	2082 2082
4.57 - 38	2408	1036	2370

Table 4. MIDDLE GUST IV Seismic Data (Ref. 16)

H	c_{p}	Cs	Υ
(m)	(m/s)	(m/s)	(kg/m ³)
0 - 3.0	335	168	2000
3.0 - 7.6	731	260	21 60

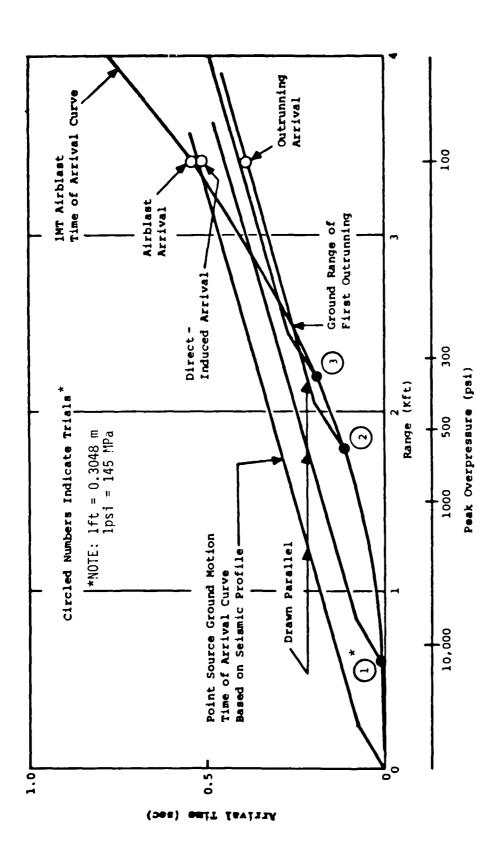
Table 5. MISERS BLUFF Seismic Data (Ref. 16)

Н	С _р	Cs	Υ
(m)	(m/s)	(m/s)	(kg/m ³)
0 - 1	335	213	1600
6.1 - 12.2 12.2 - 43	518 1707	2 60 305	1760 1920

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Table 6. PRE-DICE THROW Seismic Data (Ref. 16)

H (m)	C _p (m/s)	C _s (m/s)	r (kg/m ³)
2.4 - 33 33 - 366	1877 1981	20 4 603	17 60 1 920



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Figure **6**. Prediction of the ground shock arrival based on the site geology.

is due to the fact that upstream induced motions are not of sufficient magnitude to create very strong oscillations and, therefore, they are not observed in the data. By the time the direct-induced head waves arrive from the origin, the upstream-induced and the direct-induced waves superimpose to produce the relatively strong oscillation observed in the data.

3.3. Velocity Scaling Factor

To scale the velocity, V, of the waveform, a value, V_1 , equal to the value of the first peak of the outrunning wave was used. Figures 7 and 8 show the vertical and horizontal relations of this value with range (scaled to yield) at different depths of several experiments. (The values at different depths fall within the scatter of the data.) The equations that describe the two lines are:

$$V_1 = 2.70(10^6) R^{-2.24}$$
 for Vertical (4) $V_1 = 1.26(10^7) R^{-2.37}$ for Horizontal

where R = range in (m/MT $_{
m NE}^{1/3}$) and V $_1$ is given in m/s

3.4. Time Scaling Factor

As seen previously (Fig. 4), the frequencies of the waveforms provided by DIRECT COURSE and MIXED COMPANY were very different. DIRECT COURSE was accurately predicted by the curve developed by Sauer (Ref. 3), while MIXED COMPANY was overpredicted. To find the factors that influence the frequency, parametric studies were performed using the CAGGS computer code developed by Britt (Ref. 18). The parameters studied were 1) the depth of the overburden, H, and 2) the impedance ratio between the soil and the rock (or hard layer). As will be seen, the impedance ratio influences the frequency considerably.

The theory in CAGGS is based on the exact Cagniard elastic formulation of the wave propagation theory (Ref. 19). The Cagniard solution was extended by Britt to model the ground shock propagation and motions produced by airburst explosions. The air-earth environment was

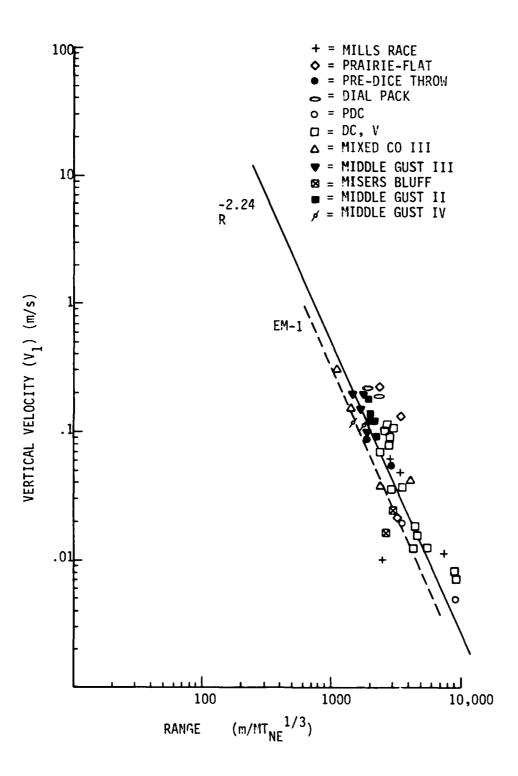


Figure 7. First peak of the outrunning wave vs. range-vertical

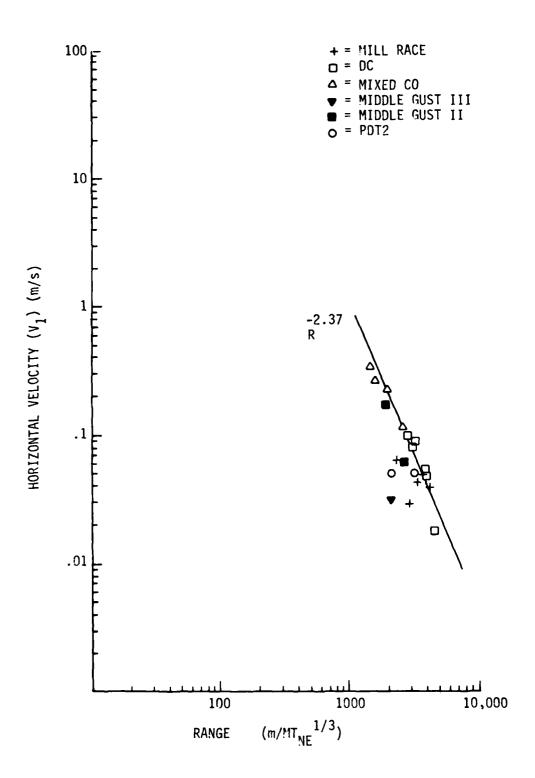


Figure 8. First peak of the outrunning wave vs. range-horizontal.

treated as three elastic layers (air, soil and rock). The explosion was replaced by a point source in an elastic fluid. The airblast was estimated empirically and was used as a source input. The CAGGS solution includes particle velocity waveforms. An example of such results is shown in Figure 9, in which the theoretical results are compared to DIRECT COURSE data at a range of 305 m. As seen in this figure, the comparison between the data and a closed form solution is good. The problem develops at large ranges where the elastic solution underestimates the magnitude of the outrunning motions considerably, and fails to estimate the frequency. CAGGS was used at small ranges for parametric studies to single out the factors that influence the outrunning wave.

Using CAGGS, the vertical and horizontal velocity of MIXED COMPANY at a range of 196 m were calculated and shown in Figure 10. As shown in Table 2, the depth to the rock in the site of this experiment is equal to 3 m. Figure 11 shows the vertical and horizontal velocity of the same profile when the depth to the rock was changed to 5.8 m. Observation of these curves reveals that the depth to the rock does not influence the shape of the horizontal velocity waveform. However, the vertical velocity shows a dependence to the depth of the softer layer. The shape of the waveform changed and the frequency was affected. The magnitude of the first peak in both the vertical and horizontal waves remained unchanged.

However, the factor that seemed to influence considerably both the vertical and the horizontal waveforms is the impedance ratio between the soil and the rock of the site. The impedance of each material was calculated using the density, ρ , and the P-wave speed, C_p , of the material. Figure 10 shows the calculated particle velocity of MIXED COMPANY at a range of 196 m and a rock velocity of 2286 m/sec (the impedance ratio is equal to 3.9). Figure 12 shows the velocity waveform of the same experiment when the P-wave speed of the rock was reduced to 1524 m/sec (the impedance ratio is equal to 2.88). When the impedance was increased to 5.78, its influence on the outrunning wave became even more noticeable (Fig. 13). A considerable change seems to occur in the frequency and general shape of both the vertical and the horizontal waves. Consequently, the impedance ratio was used, along with the characteristic time T_2 , as a scaling ratio in the derivation of a characteristic wave.

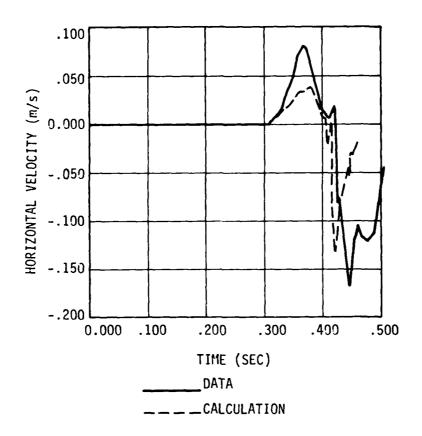
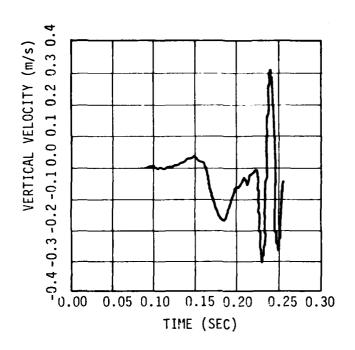
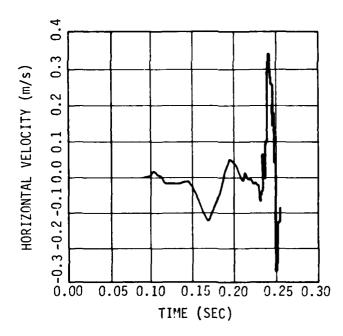


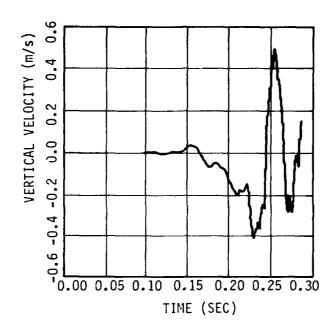
Figure 9. DIRECT COURSE data vs. calculated outrunning ground motion (range = 305 m)





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Figure 10. Calculated vertical and horizontal velocities with CAGGS--MIXED CO., range = 196 m, depth = 2.7 m, impedance ratio = 3.9



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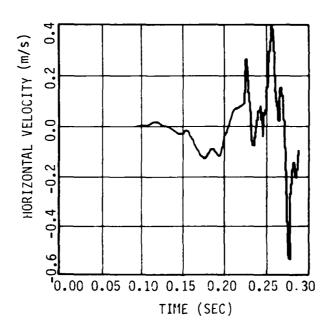
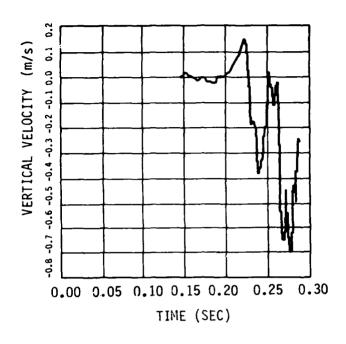


Figure 11. Calculated vertical and horizontal velocities with CAGGS--MIXED CO., range = 196 m, depth = 5.8 m



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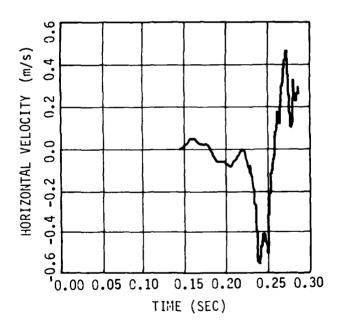
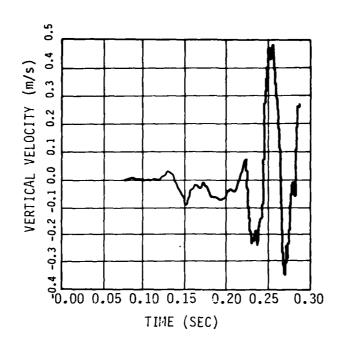


Figure 12. Calculated vertical and horizontal velocities with CAGGS--MIXED CO., range = 196 m, depth = 2.7 m, impedance ratio = 2.38



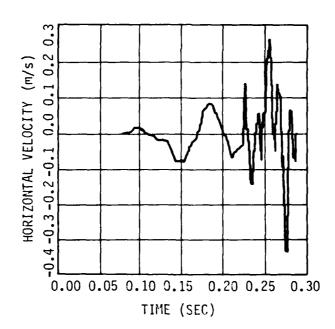


Figure 13. Calculated vertical and horizontal velocities with CAGGS--MIXED CO., range = 196 m, depth = 2.7 m, impedance ratio = 5.78

It should be noted here that the largest impedance ratio between the layers of the site should be used. This could be between the soil and the rock, the soil and the water table, or two layers of soil. At any layer interface at which the impedance ratio is calculated, the weighted average of the properties above this layer must be used.

3.5. Characteristic Wave

Using the new scaling factors, (V_1 and the impedance ratio), the outrunning waveforms from DIRECT COURSE, MILL RACE, MIXED COMPANY, MIDDLE GUST II, III, and IV, MISERS BLUFF, and Pre-DICE THROW were plotted in the normalized fashion as shown in Figure 14. A characteristic waveform was then derived by averaging this data. The vertical and horizontal waveforms that were derived are shown in Figures 15 and 16.

The characteristic waveform that was derived in this study was programmed into the Wave Synthesis Model (WSM) computer code. Several examples were prepared to show its comparison to data.

Figures 17 through 20 show the vertical and horizontal comparisons of the data with the calculations performed using WSM. The tests used in these comparisons are DIRECT COURSE, MILL RACE, MIXED COMPANY, and MIDDLE GUST II. The ranges at which the comparisons are made are at the outrunning regions; the depths are all shallow. As seen, both the magnitudes and the frequencies of these experiments are very closely predicted. The DIRECT COURSE (Fig. 17) comparison includes the entire waveform at a range of 466 m. This prediction procedure produced vertical and horizontal waveforms with an average error of 17 percent in magnitude and 5 percent in frequency. The other waveforms show the outrunning component only. MILL RACE (Fig. 18), at the range of 350 m, was predicted with an average error of 12 percent in magnitude and 12 percent in frequency. MIXED COMPANY (Fig. 19), at the range of 196 m, was predicted with an average error of 23 percent in magnitude and 8 percent in frequency. Finally, MIDDLE Gust II (Fig. 20), at the range of 107 m, had an average of 12 percent error in magnitude and 9 percent error in frequency in the predicted wave form. (NOTE: The experiments used to evaluate the procedure were those used to derive it. The procedure should be evaluated again as new outrunning data becomes available.)

A summary of the proposed prediction procedure is given in Table 7.

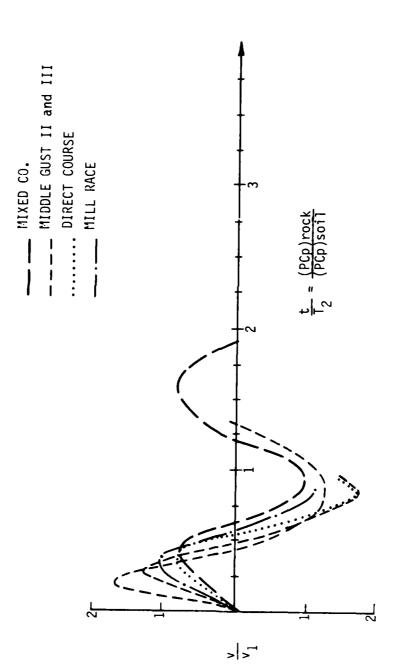


Figure 14. Different data plotted with the new scaling factors.

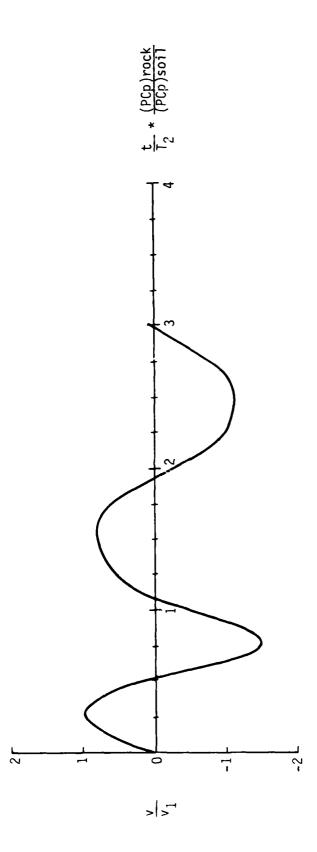


Figure 15. Characteristic vertical velocity waveform.

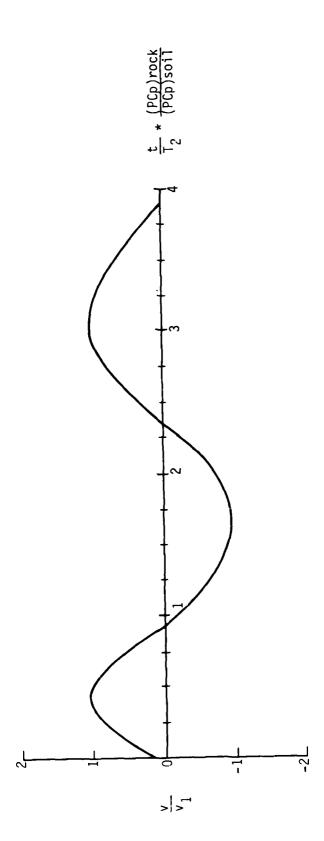
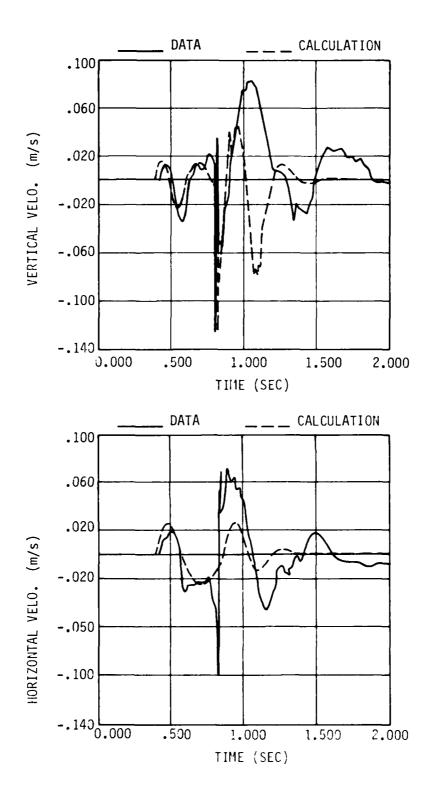


Figure 16. Characteristic horizontal velocity waveform.



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Figure 17. DIRECT COURSE data vs. calculated vertical and horizontal velocities. Range = 466 m

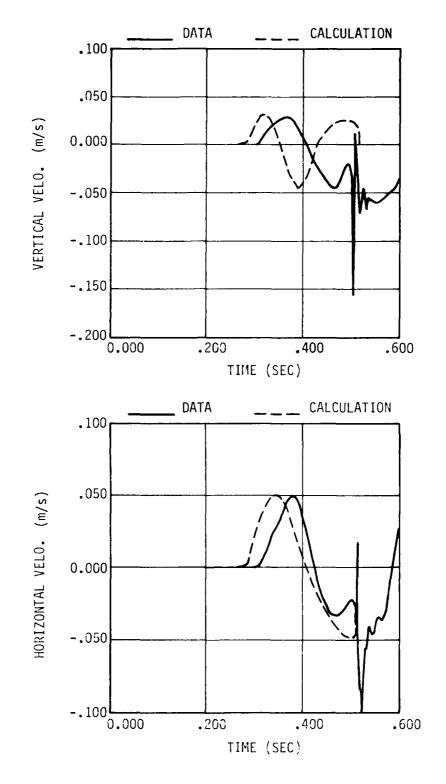
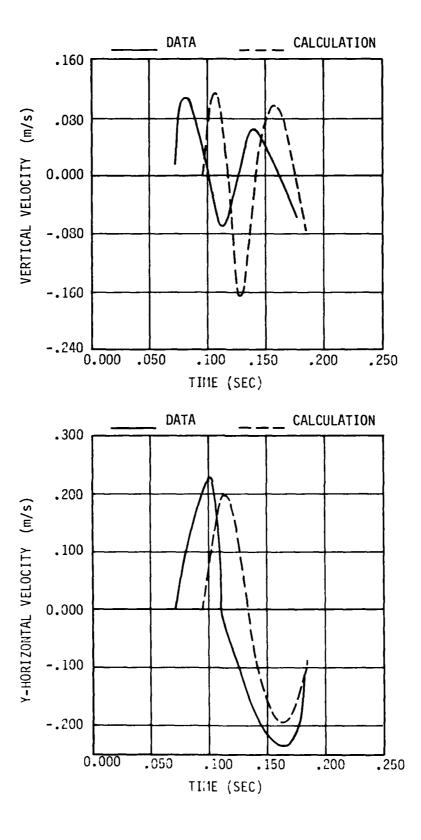


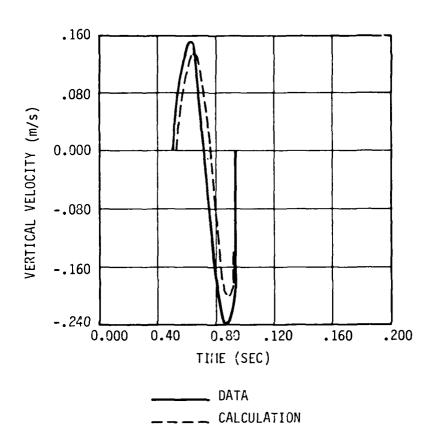
Figure 18. MILL RACE data vs. calculated outrunning vertical and horizontal velocities. Range = 350 m



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Figure 19. MIXED CO. data vs. calculated outrunning vertical and horizontal velocities.

Range = 196.6 m



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Figure 20. MIDDLE GUST II data vs. calculated outrunning vertical velocity.
Range = 106.7 m

- Table 7. Summary of the proposed prediction procedure for the outrunning waveform.
- 1) Time of arrival and duration—Time of arrival of the outrunning wave is calculated using the standard seismic velocities and layer depths of the site. In multilayered sites, it can be found using the following equation:

$$t_a^{(n)} = \left(\sum_{i=1}^{n} 2H_i \sqrt{\frac{1}{c_1^2} - \frac{1}{c_{n+1}^2}}\right) + \frac{x}{c_{n+1}} + TABO$$

where

ta(n) = arrival time of head wave from nth interface

 $H_i = thickness of the ith layer$

 C_i = compression wave velocity in the ith layer

 $C_{n+1} = compression$ wave velocity in the (n+1)th

layer

TABO = time of arrival of the airblast at the origin (for HOB explosions only)

2) Scaling factors—The velocity is scaled with a factor, V_1 , equal to the peak velocity of the first cycle in the outrunning regions. V_1 is given in m/sec and is calculated as follows:

 $V_{1}=2.70 (10^6) R-2.24$ for vertical velocities

 $V_{1}=1.26 (10^7) R-2.37$ for horizontal velocities

where: $R = \text{range of interest in m/MT}^{1/3}_{NE}$

The time is scaled with two factors:

a) a characteristic time, T_2 , where

$$T_2(sec) = 100 + 0.82 \Delta R(m)$$

in which ΔR is the difference in range between the point of interest and the point of the first outrunning.

- Table 7. Summary of the proposed prediction procedure for the outrunning waveform (Continued).
 - b) an impedance ratio, $(\rho C_p) \operatorname{rock}/(\rho C_p) \operatorname{soil}$

The time scaling factor tf is equal to

$$tf = \frac{t}{T_2} * \frac{(\rho C_p) \operatorname{rock}}{(\rho C_p) \operatorname{soil}}$$

in which t is zero at the time of arrival of the outrunning wave.

3) Waveform prediction--The waveform is predicted up to the time of arrival of the airblast. After that point other prediction procedures can be used. The following equations describe the waveforms of Figures 15 and 16:

Vertical

The state of the s

$$V_{V} = V_{1} (1) \sin \left(\frac{\pi}{.525}\right) tf \qquad \text{for } 0 < tf < .525$$

$$V_{V} = V_{1} (-1.5) \sin \left(\frac{\pi}{.55} (tf - .525)\right) \qquad \text{for } .525 < tf < 1.075$$

$$V_{V} = V_{1} (.85) \sin \left(\frac{\pi}{.85} (tf - 1.075)\right) \qquad \text{for } 1.075 < tf < 1.925$$

$$V_{V} = V_{1} (-1.1) \sin \left(\frac{\pi}{1.075} (tf - 1.925)\right) \qquad \text{for } tf > 1.925$$

Horizontal

$$V_{H} = V_{2} (1) \sin \left(\frac{\pi}{0.9}\right) tf$$
 for $0 < tf < 0.9$

$$V_{H} = V_{2} (-1) \sin \left(\frac{\pi}{1.5}\right) (tf - 0.9)$$
 for $tf > 0.9$

SECTION 4

CONCLUSIONS AND RECOMMENDATIONS

4.1. Conclusions

By the empirical analysis and the numerical calculation, several conclusions were reached.

- 1) First arrivals in the outrunning region are due to head waves from underlying layers and coincide with the P-wave arrival from the burst point.
- The first peaks of the outrunning waves attenuate at a rate of $R^{-2.24}$ for vertical and $R^{-2.37}$ for horizontal waves. Since they are produced by the headwave generated at the hard layer, they do not attenuate with depth. The magnitudes of the first peaks at depth fall within the scatter of the data with those at surface.
- 3) Height-of-burst does not influence the outrunning wave significantly.
- 4) Depth to the hard layer influences the magnitude and frequency of the vertical outrunning wave only.
- 5) The impedance ratio between the soil and rock of the site strongly influences the frequency and magnitude of the vertical and horizontal outrunning waves.
- 6) The magnitude and period of the horizontal outrunning wave are larger than those of the vertical wave

4.2. Recommendations

As data becomes available, it should be used to enrich the data base of this study and reduce the uncertainties associated with this prediction procedure.

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